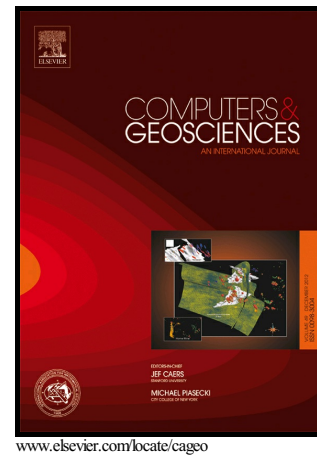


Author's Accepted Manuscript

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PII: S0098-3004(16)30445-9
DOI: <http://dx.doi.org/10.1016/j.cageo.2016.09.011>
Reference: CAGEO3842

To appear in: *Computers and Geosciences*

Received date: 3 December 2015
Revised date: 18 July 2016
Accepted date: 25 September 2016

Cite this article as: Jason Roberts, Mark Curran, Samuel Poynter, Andrew Moy, Tas van Ommen, Tessa Vance, Carly Tozer, Felicity Graham, Duncan Young, Christopher Plummer, Joel Pedro, Don Blankenship and Martin Siegert Correlation confidence limits for unevenly sampled data, *Computers and Geosciences*, <http://dx.doi.org/10.1016/j.cageo.2016.09.011>

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1 Correlation confidence limits for unevenly sampled data

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19 **Abstract**

20 Estimation of correlation with appropriate uncertainty limits for scientific
 21 data that are potentially serially correlated is a common problem made seri-
 22 ously challenging especially when data are sampled unevenly in space and/or
 23 time. Here we present a new, robust method for estimating correlation with
 24 uncertainty limits between autocorrelated series that does not require either
 25 resampling or interpolation. The technique employs the Gaussian kernel
 26 method with a bootstrapping resampling approach to derive the probability
 27 density function and resulting uncertainties. The method is validated us-
 28 ing an example from radar geophysics. Autocorrelation and error bounds
 29 are estimated for an airborne radio-echo profile of ice sheet thickness. The

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 Preprint submitted to *Computers & Geosciences*

30 computed limits are robust when withholding 10%, 20%, and 50% of data.
31 As a further example, the method is applied to two time-series of methane-
32 sulphonic acid in Antarctic ice cores from different sites. We show how the
33 method allows evaluation of the significance of correlation where the signal-
34 to-noise ratio is low and reveals that the two ice cores exhibit a significant
35 common signal.

36 *Keywords:* Unevenly sampled data, autocorrelation, bootstrapping,
37 Gaussian Kernel Method, confidence limits

38 1. Introduction

39 Sparse data correlation techniques, and the confidence limits associated
40 with them, are a keystone of quantitative analysis in geoscience. However,
41 uneven sampling of data is a common feature in many fields, and our in-
42 ability to prescribe appropriate interpolations between data may hinder the
43 statistical application of results. In many cases, this may come about as
44 an inherent sampling non-uniformity. In the case of ice cores, for example,
45 the relationship between the spatial and temporal distribution of a sample
46 material varies with depth such that uniform spatial sampling generates non-
47 uniform sampling on a temporal scale. Further difficulty arises from missing
48 data or data gaps, which may be caused by physical sample size constraints,
49 damage, or loss of samples due to contamination or analytical problems.
50 Where numerical methods require evenly sampled data, interpolation is nec-
51 essary, but must be used cautiously to avoid signal artifacts. The use of
52 common software tools to interpolate between data points often comes at
53 the expense of robustness, as bias may be introduced.

54 Rehfeld and Kurths (2014) investigated this issue in detail, bench-marking
55 a variety of techniques to overcome the challenges introduced by irregularly-
56 sampled time series. The use of a Gaussian kernel method gave a reliable and
57 robust estimation in comparison to commonly-used interpolation approaches
58 such as resampling onto a common uniform independent grid. Complications
59 arise for irregularly-sampled data with inherent autocorrelations, however,
60 as the estimation of a confidence interval, or some other measure of signifi-
61 cance, requires explicit and quantitative consideration of the autocorrelations
62 (Mudelsee, 2003; Ólafsdóttir and Mudelsee, 2014). Several methods exist for
63 the assessment of significance, for evenly sampled data, in the presence of
64 autocorrelation. Such methods include the effective spatial degrees of free-
65 dom method of Bretherton et al. (1999) which uses classical tests with a
66 reduced number of degrees of freedom to account for autocorrelations in the
67 data, and data surrogates such as bootstrapping and Fourier space methods.
68 These latter methods make no assumptions on the distribution of the data
69 (Mudelsee, 2003), so may be more appropriate for many real-world datasets.
70 Compared to standard bootstrapping techniques, Fourier space methods have
71 the advantage of preserving linear correlations, but lose many of their com-
72 putational advantages for irregularly-sampled data.

73 Here, we report the development of the Gaussian kernel method, ex-
74 tended to provide confidence interval information, with application to air-
75 borne glacier geophysical data. An evenly-sampled, highly autocorrelated
76 dataset of Antarctic ice thicknesses from the ICECAP (Investigating Cryospheric
77 Evolution through Collaborative Aerogeophysical Profiling) project (see Fig. 1
78 for location) provides a suitable test data set to validate the approach. The

79 correlation and confidence interval distribution is compared to a recently
80 published method (Ólafsdóttir and Mudelsee, 2014). Data were randomly
81 removed to simulate the effect of uneven data spacing and the resulting au-
82 tocorrelation distributions compared.

83 As a second independent demonstration of the strength of the technique
84 we compute the correlation between time series of methanesulphonic acid
85 (MSA) concentration in two Antarctic ice cores (see Fig. 1 for location).
86 MSA has been used as a proxy for Antarctic sea ice extent (Curran et al.,
87 2003), based on the production of MSA from sea ice-associated phytoplank-
88 ton which are known to be a dominant sulphur source from the sea-ice edge in
89 Antarctica (Vance et al., 2013). Confirming that a statistically significant (at
90 a 95% confidence interval) relationship exists between the two MSA records
91 supports the hypothesis that the records preserve a common environmental
92 signal.

93 While the Mudelsee (2003); Ólafsdóttir and Mudelsee (2014) method can
94 be used on unevenly spaced climate time series data, in cases where the data
95 are both unevenly spaced and on a different time base their method requires
96 interpolation or resampling. Our Gaussian Kernel-based method removes
97 the need for such resampling, making it well suited to computing correlations
98 between paleoclimate records from different locations and different archives,
99 in which different time bases are ubiquitous.

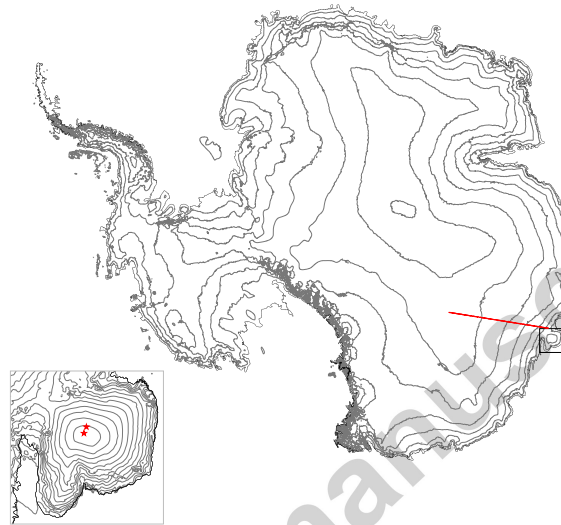


Figure 1: Location of an airborne radar transect yielding ice thickness data (red line). Elevation contours at 500 m are from Bamber et al. (2009) (grey lines) and the ice sheet grounding line is from Bindshadler et al. (2011). Inset shows the Law Dome region of East Antarctica and the sites of the two ice cores (red stars), with DSS97 being closer to the dome summit and W10K being close to the 1300 m elevation contour.

100 **2. Method**

101 *2.1. Correlation*

102 Correlations (C_{xy}) between unevenly and differently sampled series (x_i
103 and y_j) are calculated using the Gaussian kernel correlation slotting (Rehfeld
104 et al., 2011).

$$C_{xy} = \frac{1}{\sigma_x \sigma_y} \frac{\sum_{i=1}^{n_x} \sum_{j=i}^{n_y} (x_i - \bar{x})(y_j - \bar{y}) K(d_{x_i} - d_{y_j})}{\sum_{i=1}^{n_x} \sum_{j=i}^{n_y} K(d_{x_i} - d_{y_j})} \quad (1)$$

105 where the average of the two series x_i and y_j (of length n_x and n_y) is \bar{x} and
106 \bar{y} , respectively, and d_x and d_y are the independent variables (typically time
107 or distance) for x and y respectively, and may differ from each other. The
108 Gaussian kernel $K(d) = \frac{1}{\sqrt{2\pi}h} \exp(-d^2/2h^2)$ uses a width parameter (h) of
109 one quarter of the larger of the average spacing of the two data series. Unlike
110 Rehfeld et al. (2011) who normalise the signals to have zero mean and unit
111 variance, we use the original signals and correct for the mean and estimate
112 the standard deviations (σ_x and σ_y) using the same weighted summation
113 Gaussian kernel ($K(d)$) as used in Eq. 1.

114 *2.2. Bootstrapping*

115 Confidence intervals (95%) are estimated using a stationary bootstrapping
116 technique (Politis and Romano, 1994). This method accounts for per-
117 sistence (serial correlation) and the associated reduction in the effective de-
118 grees of freedom in the data (Wilks, 2006) by generating resampled data
119 series based on blocks of data from the original series. The block lengths
120 vary randomly, but the average block length is a function of the persistence
121 of the data (Mudelsee, 2003). We estimate the persistence of the data from

122 the offset required for the autocorrelation to fall to $\frac{1}{e}$, and account for the
123 uneven time intervals between data samples.

124 For each member of the dataset of bootstrap resamples, in which $n =$
125 2000, the correlation is estimated. The 95% confidence interval for the corre-
126 lation coefficient is then calculated using the bias-corrected and accelerated
127 (BCa) bootstrap method (Efron, 1987), which adjusts the result to account
128 for bias in the resampled set compared to the original point estimate of the
129 correlation coefficient from the original data.

130 Finally, in order to improve the robustness of the estimations, an addi-
131 tional 24 bootstrap resample sets of 2000 members each are generated and
132 separate estimates of the 95% confidence intervals made. The final estimate
133 of the lower (and upper) 95% confidence interval bounds is the median of
134 the 25 BCa bootstrap estimates of the lower (upper) bound. The use of the
135 median of 25 bootstrap resample sets was found to produce robust results
136 for the test datasets used here; additional resample sets may be required for
137 particularly problematic data.

138 The cumulative probability density function and its inverse, both used for
139 the BCa calculations, are based on algorithms from Abramowitz and Stegun
140 (1968) and Wichura (1988) respectively.

141 Bootstrapping methods can produce accurate estimates for relatively short
142 data series without significant autocorrelation. However, the advantages
143 of no underlying assumptions on the distribution of data comes at a cost:
144 longer data series are required when the data has significant autocorrelation.
145 Mudelsee (2003) investigated this and concluded that several hundred or
146 more data points may be required for highly autocorrelated data. Addition-

147 ally, for highly skewed data, the BCa method may predict 95% confidence
148 bounds that do not include the correlation estimate for the original data.

149 **3. Validation**

150 To validate the method, a uniformly sampled data set of 951 points was
151 used. This allowed for comparison with existing methods for uniformly sam-
152 pled data (Ólafsdóttir and Mudelsee, 2014). In addition, unevenly sampled
153 data can be simulated by removing a portion of the data on a random basis.
154 Three unevenly sampled datasets were generated, withholding 10%, 20%,
155 and 50% of the data respectively.

156 The validation data set chosen was 950 line km of ice thickness data
157 (see Fig. 2) over Eastern Antarctica (Roberts et al., 2011; Young et al.,
158 2011) (ICECAP flight ASB/JKB1a/R10Eb). This dataset covers different
159 subglacial terrains, including flat sedimentary basins (Wright et al., 2012),
160 steep-sided glacial valleys (Young et al., 2011) and rough highlands (Roberts
161 et al., 2011). This dataset has a high degree of autocorrelation at large dis-
162 placements (see Fig. 3), indicative of long characteristic length scales. The
163 high autocorrelations over relatively large spatial scales shown in Fig. 3 is
164 consistent with the results of Smith et al. (2007) who found that for East
165 Antarctica, variograms of topography could be well represented by exponen-
166 tial models with length scales up to 700 km. Since the ice sheet has a rela-
167 tively smooth surface (compared with the underlying topography), bedrock
168 topography is strongly related to ice thickness. Therefore it is not unexpected
169 that the ice thickness may show a long characteristic scale distribution as ter-
170 rain frequently exhibits long characteristic scales especially when preserved

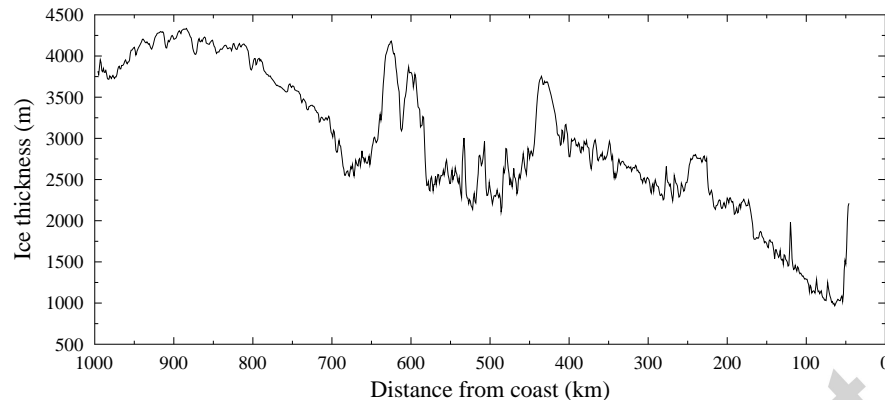


Figure 2: ICECAP radar data from flight ASB/JKB1a/R10Eb, revealing a major deep subglacial sedimentary basin, between 1000 and 700 km, in which the bed is remarkably flat; steep sided 1500 m-deep valleys between 700 and 600 km, and at 400 km; rough highland between 600 and 500 km; and flat subglacial terrain between 400 and 0 km. The profile illustrates a range of topographies measured by airborne radar profiling. The data are often hampered by losses at great ice depth, over very steep sided topography, or as a consequence of scattering at the ice surface and bed.

171 beneath ice sheets.

172 The large autocorrelation in the datasets can influence the performance of
 173 the bootstrapping estimates of the confidence interval, as the autocorrelation
 174 results in long sequences of resampled data which are very similar to the
 175 original data and, hence, artificially narrow 95% confidence intervals (e.g.,
 176 the collapsing of the 95% confidence intervals at displacements around 80
 177 and 95 km in Fig. 3). This is consistent with the finding of Mudelsee (2003)
 178 that long data series may be needed for strongly autocorrelated data.

179 Including multiple BCa bootstrap estimates and the resulting median

180 estimates for the 95% confidence interval bounds produces more robust (and
181 more smoothly varying with displacement) estimates than the Ólafsdóttir
182 and Mudelsee (2014) method.

183 The estimate is also robust to both the reduction in the amount of data
184 and unevenly sampled data, as shown in Fig. 3b and c. Results are consistent
185 even with 50% data removal. However, it should be noted that this dataset is
186 highly autocorrelated and, as such, contains much redundant information (at
187 least in terms of correlation estimates). The persistently high autocorrelation
188 at increasing displacements (Fig. 3b) is due to the non-stationary nature of
189 the ice thickness data: when the low frequency trend is removed by high-pass
190 filtering (using a Gaussian filter with an equivalent half-power width of 10
191 km) the correlations rapidly fall towards zero (Fig. 3c), although the results
192 still remain consistent with up to 50% data removal.

193 4. MSA

194 The MSA records from two distinct ice core locations (DSS97 and W10k)
195 10 km apart were compared to establish if a significant common signal exists,
196 which would support the use of MSA as a sea ice proxy in this region. For
197 broad consistency with the work of Curran et al. (2003), we low-pass filter
198 the two MSA records (Fig. 4) using a Gaussian filter with width $\sigma = 0.2994$
199 (equivalent half power width of 1 year), maintaining their unevenly sampled
200 and different time-bases. The correlation between the two MSA records is
201 0.394 [0.231–0.625], and as the confidence interval does not cross zero, we can
202 conclude that this correlation is significant at the 95% level. The results here
203 demonstrate coherence between two ice core records which experience differ-

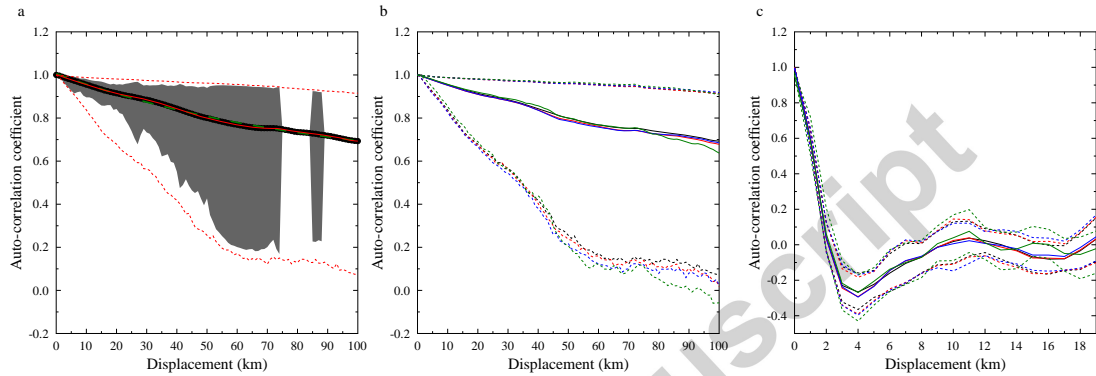


Figure 3: Autocorrelation coefficient for ICECAP ice thickness data as a function of offset displacement. a) Uniformly sampled data, red line Gaussian kernel correlation and associated 95% confidence limits (dashed red). Black line is existing method correlation calculation (Ólafsdóttir and Mudelsee, 2014) and associated 95% confidence limits (grey banding). Exponential data fit to correlation coefficient (green dashed). b) Unevenly sampled data, 0% missing (black), 10% missing (red), 20% missing (blue) and 50% missing (green). c) As per b) except for high-pass filtered ICECAP ice thickness data and different axis limits.

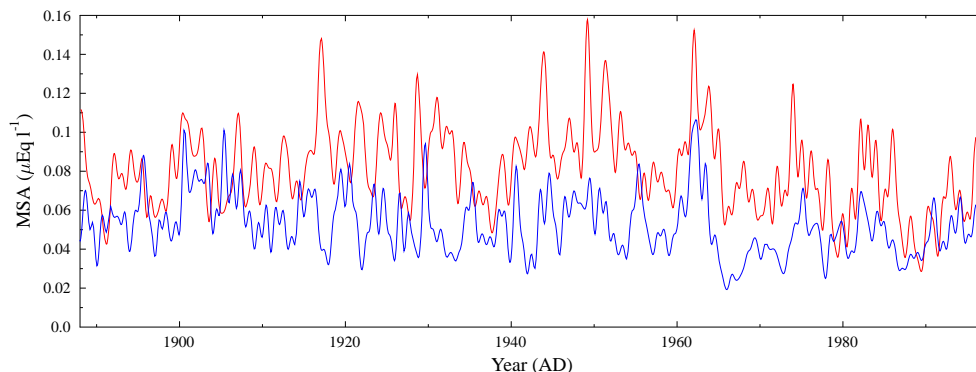


Figure 4: MSA records for two East Antarctic ice cores, DSS97 (red) and W10k (blue).

204 ent snow accumulation rates. DSS97 has a mean annual accumulation rate
 205 of 0.69 metres of ice equivalent per year (Roberts et al., 2015) and W10k
 206 of around 0.5 metres of ice equivalent per year. The different mean snow
 207 accumulation rates and associated snow densification rates result in differ-
 208 ent diffusion properties of MSA within the ice cores (Abram et al., 2008;
 209 Roberts et al., 2009), and different losses of MSA to the atmosphere, re-
 210 sulting in reduced coherence. Additionally, timing noise arises from variable
 211 snow deposition and surface relief. Despite these sources of difference, the
 212 coherent signal between the two sites provides evidence that the co-variation
 213 seen between MSA in the Law Dome region and sea-ice extent represents a
 214 regional signal that is distinct from local noise processes. This supports the
 215 interpretation of MSA as a proxy for sea-ice extent in this region.

216 **5. Software performance and concluding remarks**

217 We provide Fortran 90 source code (Intel 12.1.0, PGI 15.10-0 and gfortran
 218 4.8.4 compatible) as well as a Windows executable and MATLAB (R2015a)/Octave
 219 (3.8.1) source code in the supplementary material.

220 The bootstrap confidence interval calculation requires the calculation of
 221 50 000 (25 replicates of 2000 members) Gaussian kernel method correlations
 222 which, for larger datasets, can be quite time consuming. For example, an
 223 execution time of around 11 minutes on an Intel Core i7-4712HQ CPU for
 224 the MSA example where the sample sizes are $n_x = 1879$ and $n_y = 1491$.
 225 The majority of the time is in the nested summation ($n_x \times n_y$ repetitions)
 226 of the exponential function in Eq. 1. To decrease the execution time, the
 227 code has OpenMP compiler directives to parallelise the calculation of this
 228 nested exponential function. The speed-up obtained on the test machine is
 229 approximately 4.2 on 8 CPUs. This appears to be hardware limited, as the
 230 simultaneous execution of eight single-CPU versions of the code obtains a
 231 very similar speed-up. An OpenACC version of the code is also provided for
 232 execution on GPU or other accelerators if available.

233 The assertion that the nested exponential calculations is the most time
 234 intensive part of the calculation is confirmed by the replacement of the ex-
 235 ponential with an approximation. Using the relation

$$\exp(x) = \lim_{n \rightarrow \infty} \left(1 + \frac{x}{n}\right)^n \quad (2)$$

236 results in the following code

```
237     d=ty(j)-tx(i)
238     b=exp(-d**2/(2*h**2))/sqrt(2*pi*h)
```

239 being replaced by the approximation

```
240     d=- (ty(j)-tx(i))**2/(2.0*h**2)
241     ! if-then-else is approximation for b=exp(d)
242     if (d<-20) then
243         b=0.0
244     else
245         b=1.0+d/1024
246         b=b*b
247         b=b*b
248         b=b*b
249         b=b*b
250         b=b*b
251         b=b*b
252         b=b*b
253         b=b*b
254         b=b*b
255         b=b*b/sqrt(2*pi*h)
256     endif
```

257 This version of the code runs around 4.6 times faster than the version us-
258 ing the exponential function. Note that this version of the code is quicker (by
259 around 25%) compared to replacing all the `b=b*b` with a single `b=b**1024`.
260 This approximate version is included in the source code. For the MSA exam-
261 ple the difference in the calculated correlation value is -0.005%, with larger
262 differences in the confidence interval of around 0.8%. Since the confidence
263 interval is only an estimate, this difference may be acceptable for some ap-

264 plications with very large datasets and associated long execution times.

265 The method presented here combined with freely available software pro-
266 vides a new and valuable tool for evaluation of correlation significance in
267 common circumstances of unevenly and differently sampled, autocorrelated
268 data series.

269 **6. Acknowledgements**

270 Funding for this work was provided by the UK Natural Environment Re-
271 search Council grant NE/D003733/1, NSF grant ANT-0733025, the Jackson
272 School of Geosciences and the G. Unger Vetlesen Foundation. The Australian
273 Antarctic Division provided funding and logistical support (AAS 757, 1172,
274 2198, 3025, 3103, 4061, 4077 and 4346). This work was supported by the
275 Australian Government's Cooperative Research Centres Programme through
276 the Antarctic Climate and Ecosystems Cooperative Research Centre (ACE
277 CRC). This research was supported under Australian Research Council's
278 Special Research Initiative for Antarctic Gateway Partnership (Project ID
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