Laboratory investigation of capillary trapping under mixed-wet conditions

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[1] Remaining oil saturation established by waterflooding was measured in Indiana limestone in its original, water-wet state and under mixed-wet conditions established by adding organic acid to the oil phase. The porous plate technique was used to establish initial oil saturations ranging from $S_{nwi} = 0.23$ to 0.93 under capillary-dominated conditions. For water-wet conditions, the residual oil saturation increased linearly with its initial saturation. In contrast, the remaining oil saturation under mixed-wet conditions, S_{nw} , displayed three distinct regimes. First, S_{nw} increased with its initial saturation up to $S_{nwi} = 0.58$. Next, S_{nw} decreased from $S_{nwi} = 0.58$ to 0.76. Finally, S_{nw} increased again as S_{nwi} approached one. The nonmonotonic dependence of S_{nw} on S_{nwi} at $S_{nwi} > 0.5$ is well described by a concave-up quadratic function, and may be a salient feature of mixed-wet rocks.

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1. Introduction

[2] As a wetting phase displaces a nonwetting phase in a porous medium, a portion of the latter becomes immobilized by capillary forces in the form of discontinuous pore-scale ganglia. It is this phenomenon, known as capillary trapping, that prevents the complete extraction of nonaqueous phase liquid (NAPL) contaminants from aquifers and oil from geological reservoirs. In the context of carbon storage, capillary trapping renders a fraction of the injected CO₂ immobile against advection [Iglauer *et al.*, 2011].

[3] Capillary trapping is controlled primarily by the pore structure, wettability, and properties of the fluids and the flow. This paper considers the impact of wettability on capillary trapping in two-phase systems, focusing specifically on mixed-wet conditions. A mixed-wet state is where the porous medium exhibits pore-scale and subpore-scale heterogeneity in wettability such that parts of the pore surface are wetting to the nonaqueous phase while the rest is wetting to the aqueous phase. For example, many oil reservoirs are believed to be mixed wet [e.g., Wood *et al.*, 1991; Sutanto *et al.*, 1990; Øren *et al.*, 1998; Jerauld and Rathmell, 1997]. Here, the heterogeneity in wettability emerges from the preferential migration of oil into larger pores in reservoirs which are initially water-wetting, and the subsequent alteration of grain surfaces in contact with the oil to an oil-wetting state by, e.g., asphaltene adsorption [Kaminsky *et al.*, 1993]. Under this conceptual framework, the remaining surfaces, in the corners of pores and in smaller

pores, remain strongly water wet (see Anderson [1986] for a review). By analogy, NAPLs may render a contaminated aquifer mixed wet [Powers and Tamblin, 1995; Powers *et al.*, 1996; Dwarakanath *et al.*, 2002]. At a CO₂ storage site, exposure to CO₂ may alter the wettability of the grain surface by, e.g., dissolving the surface locally and enhancing the surface roughness at the micron-scale [Espinoza and Santamarina, 2010] or exposing a different mineral [Wang *et al.*, 2012]. Here, wettability is expected to change as the geochemical reaction progresses.

[4] It is well established that the fractional volume of nonwetting phase remaining in a water-wet rock after waterflooding (aqueous phase injection) increases monotonically with its initial saturation [e.g., Stegemeier, 1977; Land, 1968; Pentland *et al.*, 2011; Tanino and Blunt, 2012]. In contrast, the dependence of the nonwetting phase saturation on its initial saturation under mixed-wet conditions has not been studied comprehensively: existing data are focused on $S_{nwi} > 0.65$, where S_{nwi} is the fractional volume of the pore space of the rock initially occupied by the nonwetting phase [Salathiel, 1973; Graue *et al.*, 1999]. Nevertheless, it has been shown that waterflooding continues to displace oil in rocks exposed to crude oil even after 5–10 pore volumes (pv) of injection [Salathiel, 1973; Zhou *et al.*, 2000; Jadhunandan and Morrow, 1995]. This behavior is characteristic of mixed-wet conditions [Anderson, 1987] and is attributed to the oil remaining connected within the rock in the form of thin films or layers along the grain surface [Iglauer *et al.*, 2012] even at low saturations [Salathiel, 1973]. We refer to the nonwetting phase saturation in a mixed-wet system during waterflooding as *remaining* saturation, to differentiate it from the asymptotic *residual* state in which the nonwetting phase is fully disconnected.

[5] This paper presents laboratory measurements of the remaining nonwetting (nonaqueous) phase saturation in Indiana limestone in its original, water-wet state, and under mixed-wet conditions. The nonwetting phase was oil, and

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the wetting phase was brine. The porous plate technique was used to establish the full range of S_{nwi} . Mixed-wet conditions were established using a solution of 1.5 wt.% cyclohexanepentanoic acid in *n*-decane as the nonwetting phase, which has previously been shown to render calcite less water-wetting at room temperature [Wu *et al.*, 2008].

2. Materials and Methods

2.1. Rock

[6] The porous medium was Indiana limestone, selected because of its homogeneity at the core-scale (cm), strength, and commercial availability. In addition, the pore structure and flow phenomena in this rock have been studied extensively [e.g., Zhu *et al.*, 2010; Churcher *et al.*, 1991; Wardlaw and Taylor, 1976; Wardlaw *et al.*, 1987; Mair *et al.*, 1999; Hart and Wang, 1995; Tanino and Blunt, 2012; El-Maghraby and Blunt, 2013]. Indiana limestone is 99.0% calcite and 1.0% quartz (Weatherford Laboratories, East Grinstead) and is naturally water-wetting. Porosity and brine permeability were measured to be $\phi = 0.133 \pm 0.002$ (s.e., 10 samples) and $k = (2.0 \pm 0.2) \times 10^{-15} \text{ m}^2$ (s.e., 9 samples), respectively.

[7] The pore size distribution of the rock, as characterized by nuclear magnetic resonance transverse relaxation times, is reported by Tanino and Blunt [2012]. Similarly, Figure 1 presents mercury/vapor drainage capillary pressure measurements, $P_{c,Hg}$, and the pore entry size distribution derived from them previously by Tanino and Blunt [2012]. The pore entry size distribution is plotted as $r_p f(r_p)$, where r_p is the equivalent pore throat radii of the capillary pressure as calculated using the Young-Laplace equation and an assumed contact angle of $\theta_{Hg} = 130^\circ$ and interfacial tension of $\sigma_{Hg} = 485 \text{ mN m}^{-1}$ [Winslow, 1978; Good and Mikhail, 1981],

$$r_p = \frac{2\sigma_{Hg} |\cos \theta_{Hg}|}{P_{c,Hg}}, \quad (1)$$

and

$$f(r_p) = \frac{dS_w}{dr_p} \quad (2)$$

is the probability distribution function for pores with a capillary entry pressure corresponding to r_p .

2.2. Aqueous and Nonaqueous Phases

[8] The same wetting phase was used in all experiments and was prepared by mixing an aqueous solution of 5 wt.% sodium chloride and 1 wt.% potassium chloride with some crushed rock on a magnetic stirrer for a minimum of 48 h. The brine was subsequently filtered, then degassed by vacuum. For the measurement of capillary trapping under mixed-wet conditions, a solution of 1.5 wt.% cyclohexanepentanoic acid (Sigma Aldrich, 98%) in *n*-decane was used as the nonwetting phase. For unaltered (i.e., water-wet) conditions, either pure *n*-decane (Sigma Aldrich, $\geq 99\%$) or *n*-octane (Sigma Aldrich, $\geq 99\%$) was used instead.

[9] The basic properties of the liquids are presented in Table 1. The density, ρ , of all liquids were measured at 20°C (Anton Paar DMA 48 or 500). The kinematic viscos-

Figure 1. (a) Normalized drainage capillary pressure ($P_{c,Hg}/(2\sigma_{Hg} |\cos \theta_{Hg}|) (\mu\text{m}^{-1})$) required to achieve a given wetting phase saturation S_w and (b) the corresponding pore entry size distribution ($r_p f(r_p)$) as reported in Tanino and Blunt [2012]. Dashed lines depict $(S_w, P_{c,Hg}/(2\sigma_{Hg} |\cos \theta_{Hg}|)) = (0.49, 2.7 \mu\text{m}^{-1})$ [cf. equation (3)]. Dotted lines depict $S_w = 1 - S_{nwi,c} = 0.89$ and the corresponding $P_{c,Hg}/(2\sigma_{Hg} |\cos \theta_{Hg}|) = 0.41 \mu\text{m}^{-1}$ (Figure 1a) and $r_p = 2.4 \mu\text{m}$ (Figure 1b).

ity of the *n*-decane + organic acid solution was measured using a glass capillary viscometer (Rheotek Type BS/U U-tube viscometer, size A) to be $\nu = 1.20 \text{ mm}^2 \text{ s}^{-1}$; the corresponding dynamic viscosity, μ , is $\mu = \nu \rho = 0.880 \text{ mPa s}$. The interfacial tension of each oil phase with the brine, σ , was measured using the pendant drop method (DSA Easydrop, Krüss GmbH, Hamburg).

2.3. Coreflood Procedure

[10] Experiments were performed on 10 cylindrical rock samples of diameter $37.69 \pm 0.01 \text{ mm}$ and length that varied between $L = 75.66 \pm 0.02 \text{ mm}$ and $89.90 \pm 0.02 \text{ mm}$. The samples were housed in custom-made, horizontal, Hassler-type pressure cells with radial confining pressure during fluid injection.

[11] Liquids were injected using high-precision syringe pumps (Teledyne ISCO 1000D). The microscopic capillary number, $Ca = \mu U / \sigma$, and the macroscopic capillary number proposed by Hilfer and Oren [1996],

Table 1. Basic Properties of the Fluids^a

	Wetting Phase	<i>n</i> -Octane	<i>n</i> -Decane	1.5 wt.% Cyclohexanepentanoic Acid in <i>n</i> -Decane
μ (mPa s)	1.0903 ^b	0.508 ^c	0.838 ^c	0.880
ρ (kg m ⁻³)	1032.0–1059.0	702.6	729.5	732.5 ± 0.02 (s.e.)
σ (mN m ⁻¹)	–	51.64 ± 0.04 ^d	52.3 ± 0.4	27.52 ± 0.08

^aUnless otherwise noted, σ is the interfacial tension of the respective nonwetting phase with the limestone-equilibrated brine; the uncertainty is standard error of the mean.

^bFor 5.8 wt.% aqueous sodium chloride solution at 100 kPa, 293.2 K [Kestin *et al.*, 1981].

^cAt 298 K [Haynes, 2011].

^d σ with water at 293.2 ± 0.1 K [Zeppieri *et al.*, 2001].

$$\langle Ca \rangle = \frac{\mu UL/k}{(P_{b,Hg}/\sigma_{Hg})\sigma}, \quad (3)$$

where μ is the viscosity of the injected fluid, U is its volumetric flow rate per unit cross-sectional area of the rock sample, and $(S_w, P_{b,Hg}/\sigma_{Hg}) = (0.49, 3.4 \mu\text{m}^{-1})$ is the inflection point in the mercury/vapor drainage capillary pressure curve (Figure 1a, dashed) were maintained at $Ca < 10^{-6}$ and $\langle Ca \rangle \leq O(1)$ to ensure capillary-dominated conditions [Chatzis and Morrow, 1984; Hilfer and Oren, 1996].

[12] Prior to each experiment, the rock sample was saturated with brine to mimic the initial state of geological reservoirs. To do so, air was removed from the test sample by alternately purging with gaseous CO₂ and applying a vacuum. The sample was then flooded with 131 to 310 pv of degassed brine. The absolute permeability, k , was determined during this step by changing the flow rate several times and, at each flow rate, measuring the pressure drop across the length of the core once steady state was achieved. The permeability was calculated as

$$k = \mu_w \frac{dU}{d(\Delta P/L)}, \quad (4)$$

where μ_w is the viscosity of the brine and ΔP is the magnitude of the pressure difference across the core. $dU/d(\Delta P/L)$ is taken as the gradient of the line of regression of U on $\Delta P/L$. The uncertainty in k is taken as that of $dU/d(\Delta P/L)$ as estimated according to Taylor [1997, Chapter 8]. The pore volume, hence ϕ , was determined from the difference in the mass of the core between its dry state prior to the experiment and its brine-saturated state.

2.3.1. Capillary Trapping Under Mixed-Wet Conditions

[13] Each experiment consists of a sequence of fluid displacement, as described below, that mimic that associated with geological carbon storage in saline aquifers and secondary oil production. After each step, the sample was removed from the Hassler cell and weighed. The corresponding oil saturation was determined from the difference in the weight of the core from its dry state. The details of specific steps are described in Tanino and Blunt [2012] and are thus omitted here.

2.3.1.1. Primary Drainage

[14] Initial oil saturation, S_{nwi} , was established by injecting the oil phase at constant pressure from one face of the rock sample against a hydrophilic permeable disk (Weatherford Laboratories, Stavanger) placed immediately downstream of the opposite face. The disk retained the oil phase

inside the sample while allowing brine through. At equilibrium, flow ceased and, accordingly, a uniform capillary pressure, P_c , and oil saturation were established across the length of the core. Each experiment is associated with one drainage P_c , which in turn corresponds to a unique S_{nwi} . The duration of primary drainage ranged from 10 to 29 days. The maximum capillary number was constrained at $\langle Ca \rangle = 2$.

2.3.1.2. Spontaneous Imbibition

[15] After S_{nwi} was established, the core was submerged in brine in a custom-made glass Amott cell (Soham Scientific, Ely) for a minimum of 27 days to allow the brine to spontaneously imbibe into the rock. The length of the core is sufficiently short that hydrostatic pressure may be neglected, and $P_c = 0$ everywhere. Accordingly, the remaining oil saturation at the end of this step is denoted by $S_{nw}^{P_c=0}$.

2.3.1.3. Waterflooding

[16] Subsequently, the core was reinserted into a Hassler cell for waterflooding. Brine was injected into the sample at a constant volumetric flow rate of $0.100 \text{ mL min}^{-1}$, which corresponds to $U = 1.5 \times 10^{-3} \text{ mm s}^{-1}$; the macroscopic capillary number varied from $\langle Ca \rangle = 0.47$ to 1.1. At selected times, the core was removed from the Hassler cell and weighed to determine the remaining oil saturation, $S_{nw}(n)$, where n is the cumulative volume of brine injection in units of pore volume (pv). The asymptotic (large n) limit of $S_{nw}(n)$ is denoted by S_{nwr} , where the subscript r denotes residual state.

[17] The effective relative permeability of the sample to the brine after 100 pv of brine injection was estimated as

$$k_{rw}^{100pv} = \frac{\mu_w}{k} \frac{U}{\Delta P/L}, \quad (5)$$

where ΔP denotes the pressure difference across the length of the core after $n = 100$ pv at a single U .

2.3.2. Capillary Trapping Under Water-Wet Conditions

[18] The procedure could be simplified for water-wet conditions, because $S_{nw}(n)$ asymptotes after 1–2 pv of brine injection [e.g., Salathiel, 1973; Jadhunandan and Morrow, 1995; Zhou *et al.*, 2000; Mungan, 1966; Wang, 1988]. Waterflooding from a given S_{nwi} rapidly establishes a unique residual oil saturation, S_{nwr}^{ww} , that is independent of waterflood rate or cumulative volume injected as long as capillary-dominated conditions are maintained. Accordingly, S_{nwr}^{ww} was measured once or twice only, and permeability at S_{nwr}^{ww} , denoted by $k(S_{nwr}^{ww})$, was calculated using

Table 2. Summary of Coreflood Experiments Under Water-Wet Conditions, in Order of Increasing S_{nwi} ^a

Experiment	ϕ	$k (\times 10^{-15} \text{ m}^2)$	S_{nwi}	S_{nwr}^{ww}	k_{rw}^{ww}	$k_{rw}^{ww} / (1 - S_{nwr}^{ww})$
A ^b	0.131	3.12 ± 0.07	0.23	0.12 ± 0.02	0.280 ± 0.007	0.318 ± 0.008
B ^c	0.136	N/A	0.40	0.28 ± 0.03	0.178 ± 0.002	0.246 ± 0.002
C	0.121	1.6	0.76	0.38	N/A	N/A
D ^b	0.139	2.3 ± 0.1	0.86	0.47 ± 0.01	0.101 ± 0.005	0.189 ± 0.009
E	0.136	2.13 ± 0.02	0.91	0.48	N/A	N/A

^aThe uncertainty in S_{nwr}^{ww} is the difference in measured S_{nwr}^{ww} before and after the residual-state permeability measurements. The fractional uncertainty in $k_{rw}^{ww} = k(S_{nwr}^{ww})/k$ is approximated as the sum of the fractional uncertainty of k and $k(S_{nwr}^{ww})$, which in turn represent the uncertainty in the respective $dU/d(\Delta P/L)$ (equation (4)). The oil phase was *n*-octane in experiments B and C, and *n*-decane in others.

^bPreviously reported in Tanino and Blunt [2012].

^cAbsolute permeability was not measured on the core used in this experimental run. $k = (2.13 \pm 0.02) \times 10^{-15} \text{ m}^2$ of a core with a similar ϕ is used instead to calculate k_{rw}^{ww} and $\langle Ca \rangle$.

equation (4) instead of equation (5), i.e., ΔP and the corresponding U were measured at multiple flow rates. At least $n = 11$ pv of brine was injected before the first S_{nwr}^{ww} measurement was taken; the second S_{nwr}^{ww} measurement was taken after $n = 49$ – 69 pv of cumulative brine injection. Also, the residual saturation established by waterflooding is the same as that established by spontaneous imbibition under strongly water-wet conditions [Johannessen *et al.*, 2007; Zhou *et al.*, 2000]. Accordingly, $S_{nw}^{pc=0}$ was not measured explicitly.

[19] The maximum capillary number during primary drainage was constrained at $\langle Ca \rangle = 0.43$ – 0.73 . The maximum capillary number reached during waterflooding was $\langle Ca \rangle = 0.54$.

3. Results

3.1. Permeability to Brine at Remaining Oil Saturation

[20] It can be readily shown that the relative permeability of an idealized porous medium comprised of parallel, identical capillary tubes during waterflood, k_{rw} , is equal to fractional volume of the pore space that conducts brine, i.e., $k_{rw} = 1 - S_{nw}$, with the assumption that a tube can only be occupied by one phase [Dullien, 1991, Chapter 3.2]. The deviation of k_{rw} from $1 - S_{nw}$ is thus a measure of the bias in the spatial distribution of the oil phase during waterflood. For example, in strongly water-wet systems, the nonwetting phase is retained in the larger pores and the wetting phase flows through the smaller pores [Iglauer *et al.*, 2010, 2011] and, accordingly, $k_{rw} \ll 1 - S_{nw}$.

[21] The relative permeability to brine at remaining/residual oil saturation under water-wet conditions, k_{rw}^{ww} , and under mixed-wet conditions, k_{rw}^{100pv} , are presented in Tables 2 and 3, respectively. Both k_{rw}^{ww} and k_{rw}^{100pv} are smaller than one for all $S_{nw} > 0$, as expected. The small $k_{rw}^{ww} / (1 - S_{nwr}^{ww}) \leq 0.3$ observed in experiments using pure *n*-octane or *n*-decane as the oil phase is consistent with a strongly water-wetting system [Tanino and Blunt, 2012; Skauge and Ottesen, 2002; Oak *et al.*, 1990; Amaefule and Handy, 1982]. Under mixed-wet conditions, the relative permeability was larger $[k_{rw}^{100pv} / (1 - S_{nw}^{100pv}) \geq 0.5]$. An elevated relative permeability implies that brine was not restricted to the smallest pores, and thus confirms that the presence of organic acid in the oil phase has altered the

wettability of the rock. Combined, k_{rw}^{100pv} was broadly a factor of 4–5 larger than k_{rw}^{ww} .

[22] We emphasize that Darcy behavior could not be confirmed under mixed-wet conditions, because S_{nw} , hence the flow velocity for a given pressure gradient, is a function of the cumulative volume of brine injected. Thus, the permeability to brine at remaining oil saturation reported in Table 3 is strictly valid only at $n = 100$ pv of cumulative brine injection at $U = 1.5 \times 10^{-3} \text{ mm s}^{-1}$. In contrast, the flow velocity increased linearly with pressure gradient for all S_{nwr}^{ww} considered under water-wet conditions [e.g., Tanino and Blunt, 2012, Figure 8]. Accordingly, the permeability at S_{nwr}^{ww} reported in Table 2 is independent of U and n provided that capillary-dominated conditions are maintained.

3.2. Remaining/Residual Oil Saturation

[23] Here we compare the dependence of remaining oil saturation on initial saturation under mixed-wet conditions with that under water-wet conditions.

3.2.1. Water-Wet Conditions

[24] First, we consider trapping under water-wet conditions. Figure 2 presents the residual oil saturation under water-wet conditions, S_{nwr}^{ww} , measured in the present study (squares). The data coincide with measurements of $S_{nw}^{pc=0}$ established by spontaneous imbibition in Indiana limestone of similar $\phi (= 0.140)$ and $k (= 10^{-15} \text{ m}^2)$ by Wardlaw and Taylor [1976] (diamonds). The good agreement is consistent with previous observations that $S_{nw}^{pc=0} = S_{nwr}^{ww}$ under water-wet conditions (cf. section 2.3.2).

[25] In both data sets, residual saturation increases linearly with S_{nwi} over the range of S_{nwi} considered, namely $0.21 \leq S_{nwi} \leq 0.91$. A linear regression on all data shown yields a line of best-fit in the least-squares sense of (solid line)

$$S_{nwr}^{ww} = (0.06 \pm 0.02) + (0.47 \pm 0.04)S_{nwi}; \quad (6)$$

the correlation coefficient is $r = 0.980$, indicating a highly significant correlation. Note that equation (6) intersects the $S_{nwr}^{ww} = S_{nwi}$ line at a finite initial saturation of $S_{nwi} = 0.11 \pm 0.03$. Since oil saturation cannot increase during brine injection, this suggests that $S_{nwr}^{ww} \approx S_{nwi}$ (dotted line) below this critical S_{nwi} , which we denote by $S_{nwi,c}$.

[26] To understand the origin of $S_{nwi,c}$, we consider the capillary entry size distribution of the rock. From Figure 1, it can be inferred that the capillary pressure required to establish an initial nonwetting phase saturation of

Table 3. Summary of Coreflood Experiments Under Mixed-Wet Conditions, in Order of Increasing S_{nwi} ^a

Experiment	ϕ	$k (\times 10^{-15} \text{ m}^2)$	S_{nwi}	$S_{nw}^{P_c=0}$ (days)	S_{nw} (n (pv))	k_{rw}^{100pv}	$\frac{k_{rw}^{100pv}}{(1-S_{nw}^{100pv})}$	Marker in Figure 4	Best-Fit Power Function to (n (pv), S_{nw})
F	0.132	1.5	0.33	0.18 (32)	0.19 (1) 0.14 (23) 0.12 (99) 0.12 (193) 0.13 (290)	0.59	0.67	square	$S_{nw} = 0.18n^{-0.080}$
G	0.129	1.74 ± 0.01	0.58	0.48 (27)	0.32 (1) 0.28 (25) 0.25 (100) 0.24 (217) 0.24 (299)	0.37	0.49	cross	$S_{nw} = 0.32n^{-0.049}$
H	0.130	2.8	0.76	N/A	0.43 (1) 0.27 (24) 0.23 (101) 0.21 (208)	0.79	1.0	circle	$S_{nw} = 0.42n^{-0.13}$
I	0.133	1.64 ± 0.03	0.91	0.79 (273)	0.48 (2) 0.35 (25) 0.34 (100) 0.31 (216) 0.33 (299)	0.84	1.3	plus	$S_{nw} = 0.48n^{-0.078}$
J	0.139	1.46 ± 0.02	0.93	0.84 (34)	0.54 (1) 0.36 (24) 0.34 (41) 0.29 (99)	0.58	0.82	diamond	$S_{nw} = 0.53n^{-0.12}$

^aThe number in brackets under $S_{nw}^{P_c=0}$ is the duration of spontaneous imbibition at the time of $S_{nw}^{P_c=0}$ measurement.

$S_{nwi} = S_{nwi,c} = 0.11$ in this limestone is $P_c/(2\sigma\cos\theta) = 0.41 \mu\text{m}^{-1}$ (Figure 1, top). The corresponding $r_p = 2.4 \mu\text{m}$ (dotted line) falls between the first two peaks in the capillary entry size distribution, suggesting that it is a threshold between the largest, interparticle pores and the smaller, intraparticle pores. Thus, the low S_{nwi} ($< S_{nwi,c}$) regime is attributed to a state where only large macropores are invaded by the oil during primary drainage such that, during subsequent waterflooding, the brine simply flows through the smaller pores and intraparticle microporosity without expelling oil. In contrast, in the high S_{nwi} ($> S_{nwi,c}$) regime, oil invades the smaller pores and the constrictions that connect larger pores during primary drainage. This oil is partially expelled during waterflooding, resulting in $S_{nwr}^{nw} < S_{nwi}$.

3.2.2. Mixed-Wet Conditions

[27] The remaining oil saturation after spontaneous imbibition ($S_{nw}^{P_c=0}$, open circle) is compared with the residual saturation under water-wet conditions (S_{nwr}^{nw} , blue dash-dotted

line) in Figure 3a. It is evident that $S_{nw}^{P_c=0}$ increases monotonically with increasing S_{nwi} ; a linear regression to the three measurements at $S_{nwi} > 0.5$ yields

$$S_{nw}^{P_c=0} = (0.96 \pm 0.06)S_{nwi} - (0.07 \pm 0.05), \quad (7)$$

with a correlation coefficient of $r = 1.00$ (black solid line). Note that $S_{nw}^{P_c=0}$ deviates from this line at low $S_{nwi} = 0.33$ (experiment F), suggesting a regime change at $S_{nwi} \approx 0.5$; this is more apparent in the waterflood S_{nw} data, which we consider next.

[28] Consistent with previous observations (cf. section 1), the remaining oil saturation decreased continuously as waterflooding progressed at all S_{nwi} considered (Figure 4a). It is expected that, like relative permeability, the rate of oil recovery is very sensitive to the pore structure and the fluid distribution within it, which in turn varies with the evolving oil saturation. Nevertheless, the evolution in each experiment is well described by the best-fit power function in the

Table 4. Coefficients of Equation (9) Fitted to the Waterflood Data^a

Source	Rock	ϕ	$k (\times 10^{-15} \text{ m}^2)$	α_0	α_1	α_2	n (pv)	Relevant S_{nwi} Range
Present study	Indiana limestone	0.133 ± 0.002	2.0 ± 0.2	N/A	0.63 ± 0.02	-0.04 ± 0.02	0.7–1	0.58–0.93
				1.7	−2.3	1.05	23–25	
				1.4	−2.0	0.91	99–101	
				2.3	−3.2	1.31	193–217	
Salathiel [1973]	Boise sandstone	0.290 ± 0.004	1170 ± 140	2.7	−3.9	1.7	1	0.675–1.000
				2.4	−3.7	1.7	5	
				2.4	−4.0	1.9	20	

^aNote that for $n = 1$ pv in the present study, a linear function $S_{nw} = \alpha_1 S_{nwi} + \alpha_2$ was fitted instead.

Figure 2. Residual oil saturation as a function of its initial saturation under water-wet conditions (S_{nwr}^{ww} , squares). Vertical bars depict the difference between two S_{nwr}^{ww} measurements taken before and after the residual-state permeability measurements; S_{nwr}^{ww} was measured only once in experiments C[(S_{nwi}, S_{nwr}^{ww}) = (0.76, 0.38)] and E[(0.91, 0.48)]. Also shown is the remaining mercury saturation measured in Indiana limestone by Wardlaw and Taylor [1976] (mercury/vapor, diamonds). The solid line is equation (6), the line of best-fit to all data shown.

least-squares sense to the respective data (solid line); the corresponding equations are provided in Table 3. The exponent in the best-fit functions ranges from -0.078 to -0.13 . Similarly, the remaining oil in a mixed-wet sandstone during waterflooding documented by Salathiel [1973] varies like $S_{nw} \propto n^{-0.18}$ over $n = 2$ to 4600 pv (dotted line).

[29] By extrapolating the best-fit functions to lower n (early times), we can estimate the critical cumulative waterflood volume at which S_{nw} for a given S_{nwi} first falls below the equivalent residual saturation under water-wet conditions; the latter is given by S_{nwr}^{ww} interpolated from equation (6) for that S_{nwi} . The critical n ranges from $O(0.1)$ to $O(1)$, with the transition occurring at $n = 0.17$ pv for $S_{nwi} = 0.33$ and $n = 1.6$ pv for $S_{nwi} = 0.93$ (Figure 4b).

[30] We now consider the dependence of $S_{nw}(n)$ on S_{nwi} . $S_{nw}(n)$ at selected cumulative waterflood volumes are presented in Figure 3a. At $n \approx 1$ pv, S_{nw} increases monotonically with S_{nwi} (green). The data at $S_{nwi} > 0.5$ is well described by the best-fit line (solid green line):

$$S_{nw} = (0.63 \pm 0.02)S_{nwi} - (0.04 \pm 0.02); \quad (8)$$

Equation (8) underestimates S_{nw} at $S_{nwi} = 0.33$. At larger n considered, S_{nw} is clearly not monotonic. Specifically, S_{nw} increases as S_{nwi} increases from 0.33 to 0.58, then decreases slightly, then increases again from $S_{nwi} = 0.76$ to 0.93. Furthermore, S_{nw} above $S_{nwi} = 0.5$ is well described by a quadratic function of the form

$$S_{nw}(n) = \alpha_0 S_{nwi}^2 + \alpha_1 S_{nwi} + \alpha_2, \quad (9)$$

where coefficients α_0 , α_1 , and α_2 are fitted to the data in the least-squares sense (solid lines). Note that S_{nw} cannot exceed S_{nw} at earlier times, i.e., equation (9) cannot intersect equation (8). Accordingly, the concave-up profile described by equation (9) is not valid below some threshold $S_{nwi} \approx 0.5$.

[31] The only existing data of waterflood oil recovery in mixed-wet rock as a function of initial saturation and n , in which the initial saturation is established by a fluid displacement sequence representative of those that occur in the field, in the literature are those of Salathiel [1973]. S_{nw} in a Boise sandstone reported by Salathiel [1973] also displays a quadratic, concave-up dependence at high $S_{nwi} \geq 0.68$ under mixed-wet conditions (Figure 3b, dashed line). This is particularly interesting because Boise sandstone and Indiana limestone display distinctly different $S_{nwr} - S_{nwi}$ trends under water-wet conditions. Specifically, Indiana limestone displays a bilinear profile (section 3.2.1), while Boise sandstone displays the classic Land [1968]-type profile, in which the rate of increase of S_{nwr} decreases

Figure 3. Remaining oil saturation as a function of its initial saturation under mixed-wet conditions: $S_{nw}^{Pc=0}$ (open circles) and $S_{nw}(n)$ for selected $n = 0.7 - 1$ pv (green circles), 23–25 pv (cyan circles), 99–101 pv (red circles), and 193–217 pv (yellow circles). Solid lines are the line of best-fit to $S_{nw}^{Pc=0}$ (black) and $S_{nw}(1$ pv) or the best-fit second-order polynomial to $S_{nw}(n > 20$ pv), for $S_{nwi} > 0.5$; see Table 4 for numerical values of the coefficients. The dash-dotted line in Figure 3a is equation (6), and depicts the residual saturation under water-wet conditions. Also shown in Figure 3b are $S_{nw}(n)$ measured at $n = 1$ pv (crosses), 5 pv (pluses), and 20 pv (dots) in a mixed-wet sandstone by Salathiel [1973]; dashed lines are best-fit second-order polynomials (cf. Table 4). The data point for $(S_{nw}, n) = (0.19, 1$ pv) at $S_{nwi} = 0.33$ overlaps with $S_{nw}^{Pc=0} = 0.18$ and can hardly be discerned in Figure 3a.

Figure 4. The evolution of the remaining oil saturation during waterflooding for $S_{nwi} = 0.33$ (squares), 0.58 (crosses), 0.76 (circles), 0.91 (pluses), 0.93 (diamonds) (a) as S_{nw} and (b) in normalized form S_{nw}/S_{nw}^{ww} , where S_{nw}^{ww} is that predicted by equation (6). Solid lines are best-fit power functions (see Table 3). Also plotted in Figure 4a are $S_{nw}(n)$ in a mixed-wet sandstone reported in Salathiel [1973] (dots) and the corresponding best-fit power function ($S_{nw} = 0.425n^{-0.181}$, dotted line).

continuously with increasing S_{nwi} [Land, 1968]. Combined, these observations suggest that the concave-up quadratic $S_{nw}(S_{nwi})$ profile may be a characteristic feature of mixed wettability that is conserved across a wide range of pore structure.

[32] The concave-up profile may be explained qualitatively as follows. As S_{nwi} increases, an increasing number of pores and fractional surface area are modified to oil-wetting. As a result, the connectivity of the oil during waterflooding in the form of films (or layers) along the oil-wet patches of the grain surface is enhanced, offsetting the increase in S_{nwi} such that S_{nw} decreases with increasing S_{nwi} . As S_{nwi} increases further, increasingly smaller pores and corners of larger pores become oil wet. However, because the capillary pressure distribution drives the injected brine through the centers of the larger pores, the additional oil in the smaller pores is largely retained during waterflooding and S_{nw} increases with S_{nwi} .

[33] While this explanation appears plausible, a predictive quantification of this behavior poses a challenge for pore-scale modeling. Typically, pore network models incorporate wettability alteration by assigning grain surfaces in contact with the oil after primary drainage oil-wet characteristics. Then, by using the relative permeability predicted by the simulator, the evolution of the remaining oil saturation during waterflood may be predicted from a Buckley-Leverett analysis. Predictions vary considerably even for the same rock. For example, Valvatne and Blunt [2004]’s simulator yields a concave-up profile at the highest S_{nwi} (≥ 0.69) in a pore network representative of Berea sandstone. In contrast, a more recently developed simulator predicts a monotonically increasing dependence [Ryazanov et al., 2009].

[34] Similarly, predictions for the endpoint (long-time) residual saturation display qualitatively different behavior: many scenarios yield a two-regime, concave-down dependence over the full range of S_{nwi} [e.g., Blunt, 1997; Jackson et al., 2005], while others yield a constant saturation at high S_{nwi} (≥ 0.8) [Ryazanov et al., 2009]. Interestingly, Blunt [1997] did predict a three-regime trend of residual oil saturation, an increase, then a decrease, and a more gradual increase, reminiscent of the present data, as the fraction of oil-wet pores in the network increased at a specific $S_{nwi} = 0.89$. This was explained through the subtle interplay of oil layer connectivity, the connectivity of pores that remain water-filled after primary drainage, and the competing pore-scale displacement mechanisms. However, the same trend was not observed with increasing S_{nwi} .

[35] A close inspection of these studies suggests that it is indeed differences in the assigned contact angle distribution and its history, rather than the pore structure or topology, that give rise to the different trapping profiles. It is possible that accurately modeling the wettability alteration during the coreflood experiment, in a network representative of Indiana limestone, could reproduce the dependence on S_{nwi} observed in the laboratory. This is a topic for future work using more sophisticated pore network models.

4. Conclusions

[36] The impact of wettability on capillary trapping was studied explicitly by measuring the remaining oil saturation established by waterflooding in the same rock under uniformly water-wet conditions and under mixed-wet conditions. Under water-wet conditions, the residual oil saturation increased linearly with its initial saturation. In contrast, under mixed-wet conditions, the remaining oil saturation displayed a non monotonic profile, with S_{nw} increasing, decreasing, then increasing again as S_{nwi} increases from 0 to 1. S_{nw} at $S_{nwi} > 0.5$ under mixed-wet conditions is accurately described by a concave-up quadratic function.

[37] Returning to our original motivation for this study, we now consider implications for EOR, in which waterflooding is performed to reduce S_{nw} at $n \leq O(1)$ pv. In a fractured reservoir, waterflooding drives flood water through the high permeability fractures, and oil remaining in the rock matrix is expelled predominantly by spontaneous imbibition of water from the fractures. Given $S_{nw}^{P_{c=0}} \geq S_{nw}^{ww}$ (Figure 3a), mixed-wet conditions are clearly unfavorable to recovery in such reservoirs. In contrast, in reservoirs without fractures, the remaining oil is displaced directly by the flood water. The present data suggest that at low S_{nwi} , mixed-wet conditions yield higher recovery than water-wet conditions, i.e., $S_{nw} < S_{nw}^{ww}$, even at $n < 1$ pv (Figure 4b). This in turn suggests that the classic EOR technique of chemical flooding to revert the wettability of an oil reservoir to a more water-wetting state [Leach et al., 1962] may be detrimental in reservoirs where the initial (post-primary drainage) oil saturation was relatively low, e.g., reservoirs with a thick transition zone above the oil/water contact.

[38] One interesting question that emerges from the present study is the extent to which the concave-up quadratic correlation between remaining oil saturation and its

initial saturation applies to different porous media. Most water-wet rocks can be classified into two categories based on their initial-residual saturation profile: the classic, Land [1968]-type nonlinear profile and a bilinear profile. The present study indicates that rocks in different categories under water-wet conditions can display similar profiles under mixed-wet conditions. Further insights require an extension of the present laboratory investigation to a suite of porous media representing a wide range of pore-scale and subpore-scale structure.

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